

US-Japan Collaborative Research on Model Assessment of the March 11 2011 Earthquake, Tsunami Inundation and Initial Spread of Cs-137 in the Coastal Ocean

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http://fvcom.smast.umassd.edu/research_projects/FVCOM_Tsunami/index.html

Sea of Okhotsk

M9 (M8.8) earthquake center 5:46 GMT (tocal time: 14.46 PM) March 11, 2011 Depth: 10km; Location: 38.03N, 143.15E

North Korea 편양숫Pyongyang Seoul East Sea Sea of Japan

Yellow Sea West Sea South Korea

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東<u>京</u>日 和 Tokyo

M6.3-00:05:03 GMT June 3, 2011 Depth: 31 km

Eastern China Sea

☆Taipei

北京市

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Tokyo



North of Sendai (仙台)





Sendai (仙台)



Sendai (仙台)



Sendai (仙台)





Fukushima Daiichi Nuclear Plant



Fukushima Dai-ichi Radionuclides

Cs-137: Half life is 30 years Cs-134: Half life is 2 years I-131: Half life is 8 days

- Fukushima Source
- coastal water (5-10 km)
- 30 km offshore
- Japan 1960-2010 baseline
- Baltic Sea 1986 Chernobyl
- Black Sea 1986 Chernobyl
- US EPA drinking water standard







Bq is a measuring unit for radioactivity,1Bq=1 atomic disintegration per second

Radioactive Water Loading

Time	Types	Source	Amount (ton)
7:03 PM April 4 to 5:40 PM, April 10	Dumping (LLRW)	A facility at crippled nuclear power complex to make room for the storage of highly contaminated water	9070
9:00 PM April 4 to 6:52 PM, April 9	Dumping (LLRW)	A sub-drain pit of Unit 5	95 0
9:00 PM April 4 to 6:52 PM, April 9	Dumping (LLRW)	A sub-drain pit of Unit 6	37 3
?? April 1 to 5:38 AM April 6	Leaking (HLRW)	Unit 2	520 Cs-137: 9.4×10 ¹⁴ becquerels (Bq) Cs-134: 9.4×10 ¹⁴ Bq I-131: 2.8×10 ¹⁵ Bq
2:00 AM May 10 to 6:45 PM, May 11	Leaking (HLRW)	Unit 3	250 Cs-137: 9.8×10 ¹² Bq Cs-134: 9.3×10 ¹² Bq I-131: 8.5×10 ¹¹ Bq

LLRW: Low-level radioactive water; HLRW: High-level radioactive water

Questions:

- 1. Tsunami Inundation: How does Tsunami-solitary wave propagate towards the coast? In particularly, how does the water flood onto the coastal land and into Fukushima Daiichi Nuclear Plant?
- 2. How the high-radioactive water flow out onto the shelf from Fukushima Daiichi Nuclear Plant? How high percent of Cs-134 and Cs-137 could be advected to the Pacific Ocean?
- 3. If Cs-134 and Cs-137 is advected to the Pacific Ocean, how long could they be transferred to the US coast and other coasts of surrounding countries?



Fukushima Dai-ichi Nuclear Plant Facility



The wind velocity vectors at the 10-m height off Sendai, Japan



What do we need to address listed questions?

- 1. A sound earthquake model capable of reproducing the sea floor change caused by the earthquake?
- 2. A global-regional-coastal model system to resolve a) global ocean and regional circulation and b) coastal inundation. This model must includes a complete physics: realistic meteorological forcing, tides, and wave, etc.
- 3. High-resolution bathymetric data for both ocean and lands.
- 4. Detailed field measurements

In addition:

To resolve the geometry of Fukushima Daiichi Nuclear Plant, we need a resolution of ~ 5 m.







Approach

Step 1:

Run Global-FVCOM with the hindcast meteorological forcing, river discharges and tides (8 major tidal constituents) from January 1, 2011 to March 15, 2011. The experiments are made with inclusion of daily SSH and SST data assimilation.

Step 2:

Re-run Global-FVCOM with a restart time at 00:00 GMT, March 11, 2011 and add the seafloor change at 5:46 GMT (the time when the earthquake occurred). A total of 6 earthquake models are considered.

Step 3:

Run the high-resolution FVCOM coastal inundation model (with a resolution up to 5 m) with the same meteorological forcing from Global-FVCOM and initial earthquake-induced sea level setup from the earthquake model.

Step 4:

Release tracers and track them in a 3-D high-resolution field.



Calculated bottom elevation and depression change caused by March 11 M=9 and M=7.9 quakes estimated by five seismic models:

Case-a: Wei and Sladen (2011)

(with data sources from teleseismic P, SH and long period surface waves plus static GPS);

Case-b: Fujii and Satake (2011)

(with data sources from DART tsunamigrams and tidal guages);

Case-c: Shao et al (2011)

(with data sources from the same data source as *Wei and Slanden*'s except no GPS);

Case-d: *Pollitz and Burgmann* (2011) (with data sources from static GPS and smoothed from 15,876 patches);

Case-e: Yagi (2011) (with data sources from teleseismic body waves).







The propagation speed of the Tsunami Solitary Wave

Red line:

The location of the wave crest.

Number:

Propagation speed (m/s)

Tsunami-derived water levels

Photos of strong vortex during 2011 event provided by Kyeong Ok Kim at KORDI, South Korea.

Survey regions of the Tsunami inundation area by Japanese survey teams.

Total: 16 regions

Types	Inundated area (km²)
Measured	32.89
Case-a	26.99
Case-b	28.94
Case-c	30.94
Case-d	26.87
Case-e	20.39

Modeled: case-b

Modeled: case-e

Questions:

- How could we track the CS-137 spreading?
- How are the results sensitive to model resolution?

The movement of the center of the dye patch is controlled by

$$\frac{dx_c(t)}{dt} = \frac{1}{\hat{C}} \left[\int \int_A \bar{C}(\bar{u}h) dx dy + \int \int_A h \overline{c' u'} dx dy + \int \int_A h \overline{c' u'} dx dy + \int \int_A x h \bar{C}_o dx dy \right]$$

$$\frac{dy_c(t)}{dt} = \frac{1}{\hat{C}} \left[\int \int_A \bar{C}(\bar{v}h) dx dy + \int \int_A h \overline{c' v'} dx dy + \int \int_A h \overline{c' v'} dx dy + \int \int_A y h \bar{C}_o dx dy \right]$$

Chen et al. (2008): A model-dye comparison experiment in the tidal mixing front zone on the southern flank of Georges Bank, JGR-Ocean, 113, C02005

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Cs-137 Spreading Experiments

The WHOI Field Measurement Program in June 2011

Model-data comparisons

Cs-137 Distribution in June, 2011

Time series of the Cs-137 concentration in the sediment over the shelf.

Time series of the Cs-137 concentration in the sediment at near-shore stations

Do we need to include the non-hydrostatic dynamics?

~1 ~1000 km

Hydrostatic

 $\frac{\text{Surface elevation}}{\text{Local water depth}} = \frac{\zeta}{H} << 1$

Large-scale motion in which the vertical motion is at least one order of magnitude smaller than the horizontal motion.

Vertical convection, over-turning, and high frequency internal waves are not resolved. ~ a few meters

Non-hydrostatic

 $\frac{\text{Surface elevation}}{\text{Local water depth}} = \frac{\zeta}{H} \sim 1$

Small-scale motion in which vertical motion is the same order of the horizontal motion.

Vertical convection, over-turning, and high frequency internal waves can be resolved.

Hydrostatic

Non-Hydrostatic

Red: Non-hydrostatic

Black: Hydrostatic

Black: Hydrostatic

Summary

- 1. The global-regional-coastal FVCOM system is suitable for the Tsunami inundation simulation;
- The impact of radionuclides on the Pacific Ocean needs to be assessed with attentions to the multi-scale ocean processes along the Japan's coast;

Fixed depth@0 m

http://marinedebris.noaa.gov/info/patch.html#lit

Fixed depth@ 50 m

