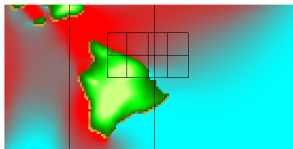
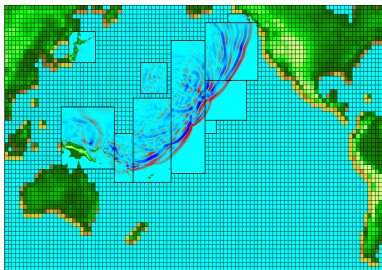


GeoClaw and Tsunami Modeling: unique numerical aspects and future directions

David L. George¹

¹Cascades Volcano Observatory, U.S. Geological Survey

PASI, Valparaiso, Chile, Jan. 2013



Acknowledgments and Collaborators

Numerical methods and software development:

- Randall LeVeque, University of Washington.
- Marsha Berger, Courant Institute, NYU.
- Kyle Mandli, University of Texas, Austin.
- Donna Calhoun, Boise State University.
- David Ketcheson, KAUST.
- Dave Yuen, University of Minnesota.

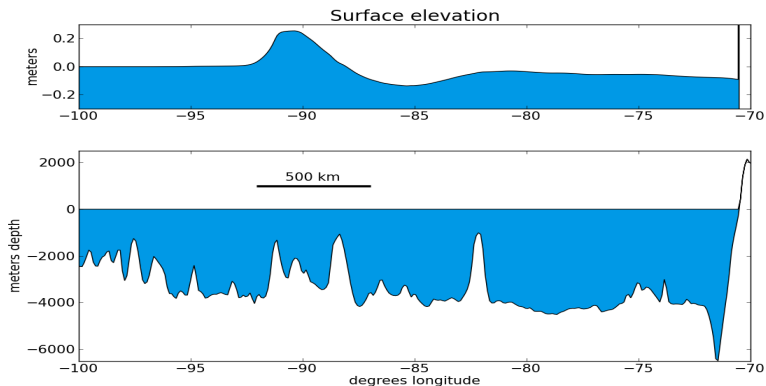
Geophysical modeling:

- Richard Iverson, U.S. Geological Survey.
- Roger Denlinger, U.S. Geological Survey.

- Survey of some unique numerical issues for tsunami modeling
 - well-balancing for ocean propagation
 - depth-positivity (wet/dry problem)
 - shoreline instabilities
 - ill-posedness and instabilities (*e.g.*, resonance, roll waves, etc.)
- Introduction to granular-fluid flows
 - modeling multiphase “non-rheological” flows
 - coupling with tsunamis?

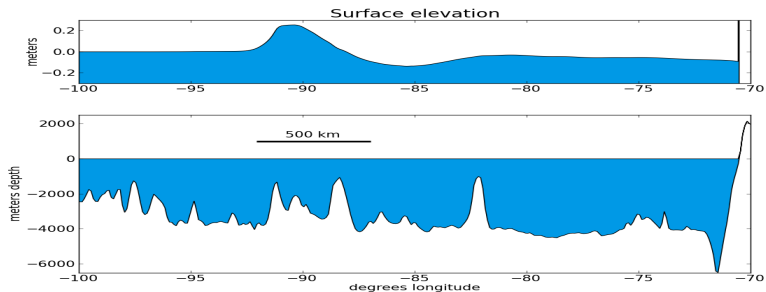
Steady-states and well balancing

- well-balancing has to do with numerically maintaining particular steady-states
- required to resolve small perturbations to those steady-states
- cannot be determined by order of accuracy or convergence
→ should be satisfied on coarse grids



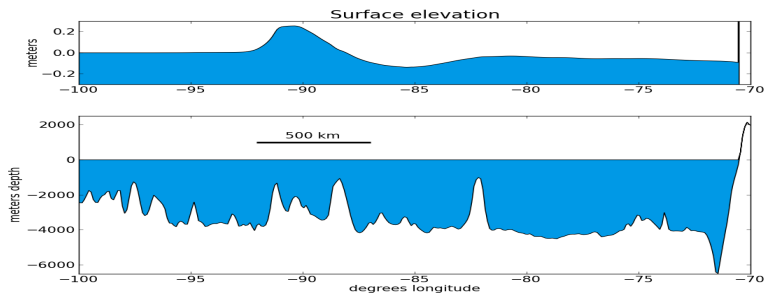
Steady-states and well balancing

$$\frac{\partial h}{\partial t} + \frac{\partial(hu)}{\partial x} + \frac{\partial(hv)}{\partial y} = 0,$$
$$\frac{\partial(hu)}{\partial t} + \frac{\partial}{\partial x}(hu^2 + \frac{1}{2}gh^2) + \frac{\partial(huv)}{\partial y} = -gh \frac{\partial b}{\partial x},$$
$$\frac{\partial(hv)}{\partial t} + \frac{\partial(huv)}{\partial x} + \frac{\partial}{\partial y}(hv^2 + \frac{1}{2}gh^2) = -gh \frac{\partial b}{\partial y}.$$



Steady-states and well balancing

$$\frac{\partial h}{\partial t} + \frac{\partial(hu)}{\partial x} = 0,$$
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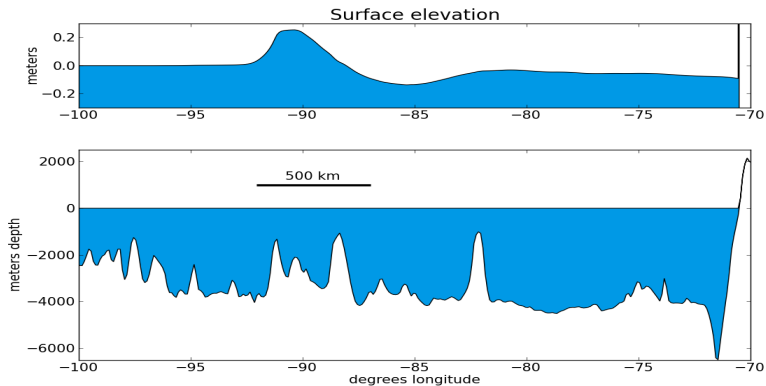


Steady-states and well balancing

Flow dynamics are small perturbations to a **balanced steady state**:

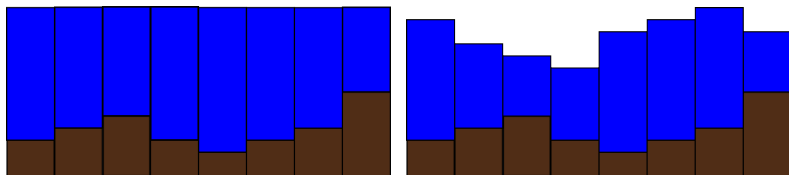
$$q_t + \mathcal{A}(q)q_x = \psi(q, x, y),$$

$$\mathcal{A}(q)q_x \approx \psi(q, x, y)$$



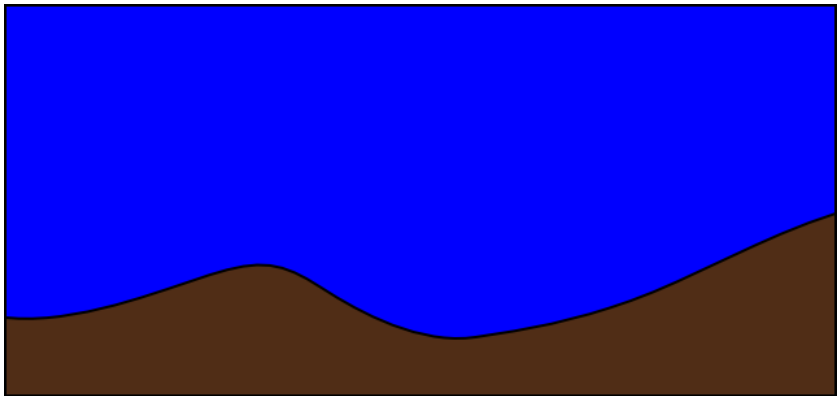
Steady-states and well balancing

- applications with delicate balanced steady states require specialized methods.
- a numerical scheme is **well-balanced** if it exactly preserves discontinuous steady-state weak solutions.
- traditional schemes do not satisfy this property.



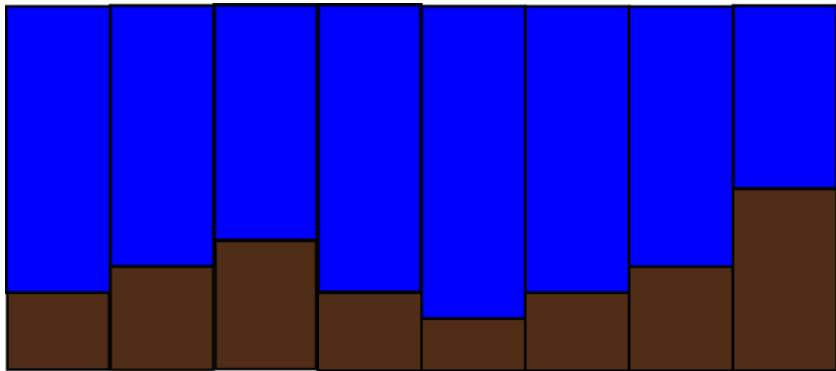
Balance law $q_t + f(q)_x = \psi(q, x)$

- Steady states arise when $f(q)_x = \psi(q, x)$



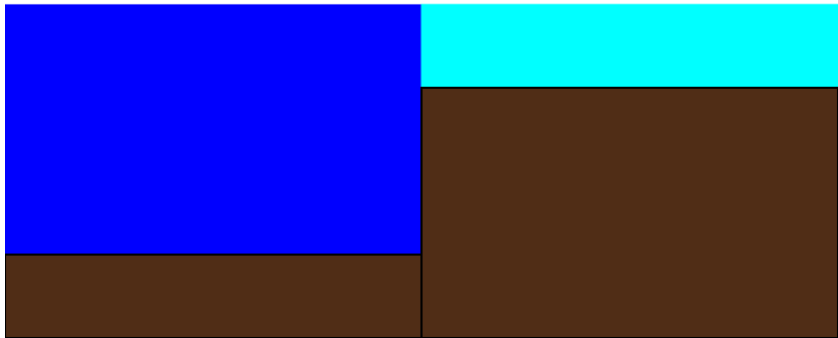
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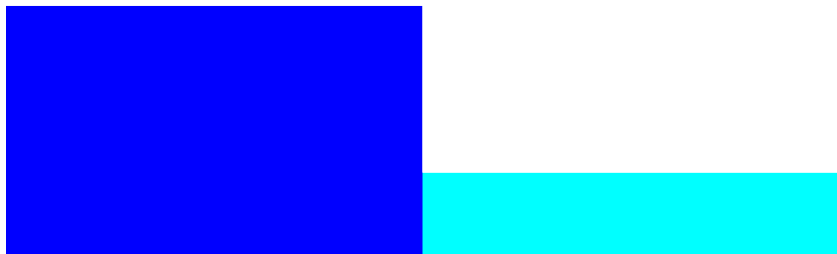
Balance law $q_t + f(q)_x = \psi(q, x)$: fractional step method

- step 1: $Q^n \rightarrow Q^*$: solve homogeneous problem:
 $q_t + f(q)_x = 0$.



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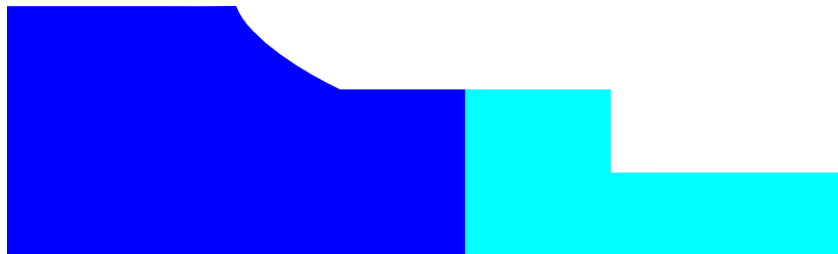
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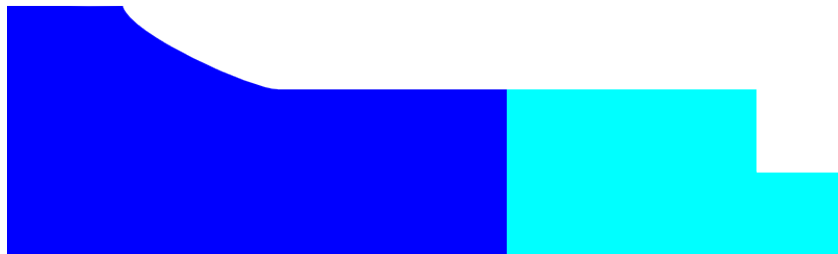
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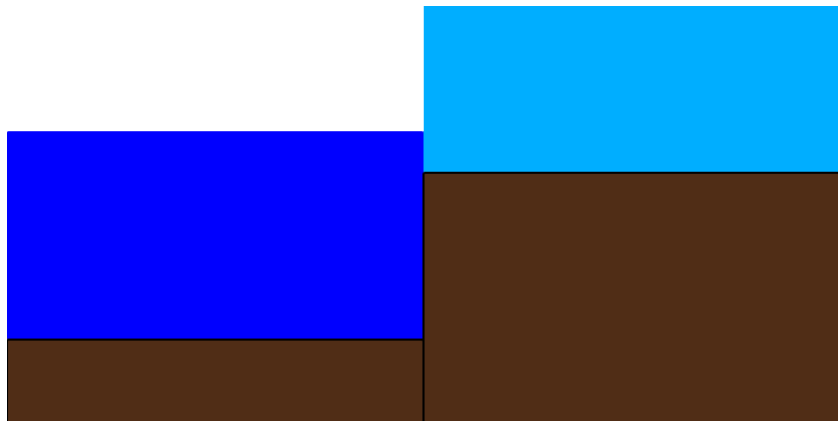
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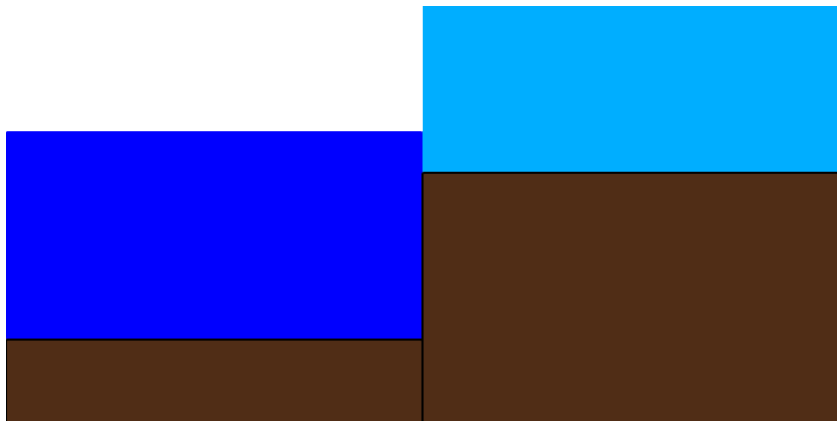
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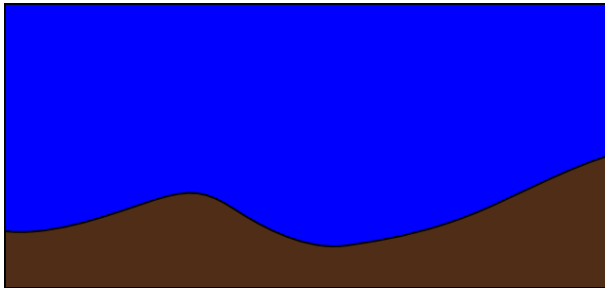
- step 2: $Q^* \rightarrow Q^{n+1}$: solve ODEs for the source term:
 $q_t = \psi(q, x)$



Balancing steady states: $q_t + f(q)_x = \psi(q, x)$

- Preserving steady states $q_t + f(q)_x = \psi(q, x)$.
- $\mathcal{W} = [q, \varphi(q), b]^T$: $\mathcal{W}_t + \mathcal{A}(\mathcal{W})\mathcal{W}_x = 0$.
- The steady state field

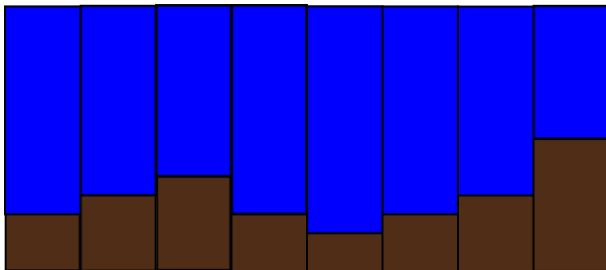
$$\mathcal{A}(\mathcal{W})\mathcal{W}_x = 0, \quad \Rightarrow \mathcal{W}_x = r^0(\mathcal{W}), \quad \lambda^0(\mathcal{W}) = 0$$



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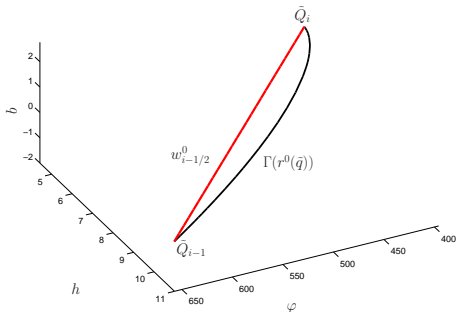
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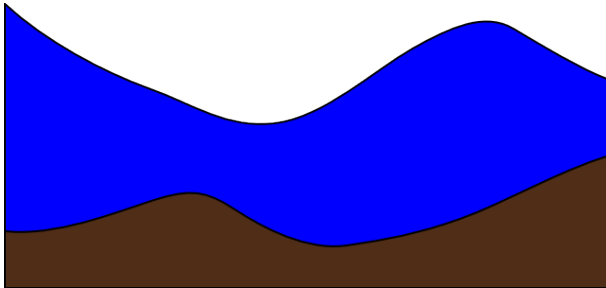
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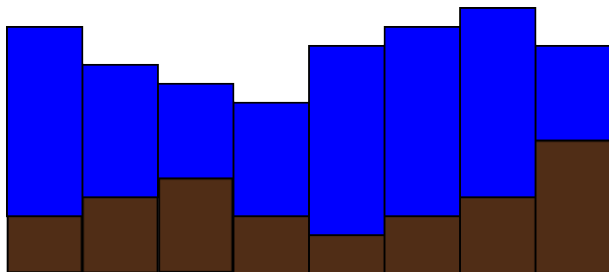
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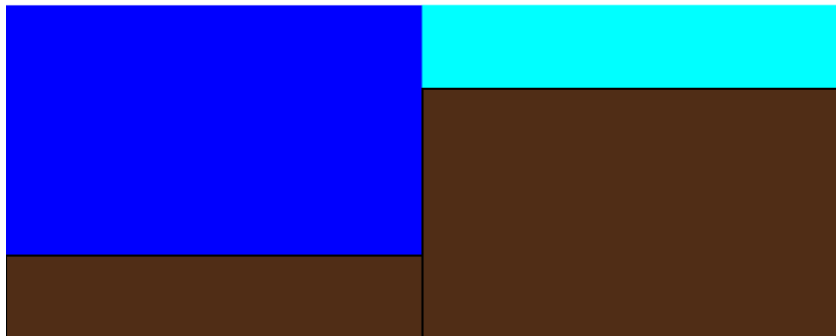
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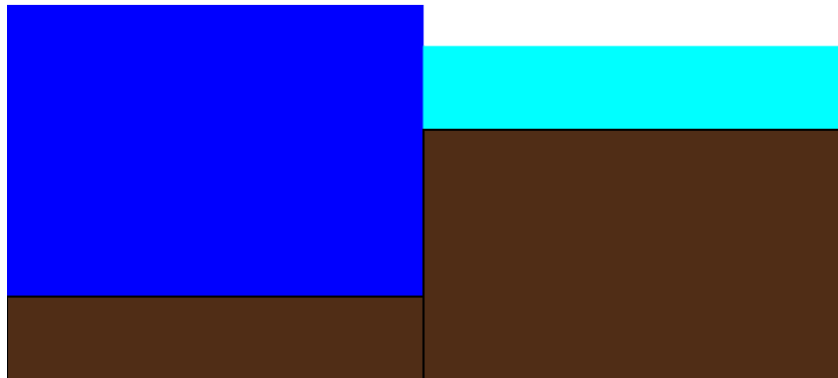
Balance law $q_t + f(q)_x = \psi(q, x)$: steady state wave

- solve augmented homogeneous system: $\mathcal{W}_t + \mathcal{A}(\mathcal{W})\mathcal{W}_x = 0$.
- motionless steady states: stationary steady state wave only



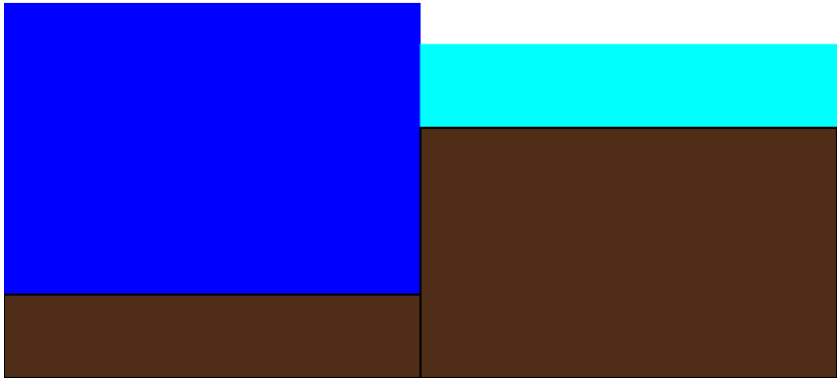
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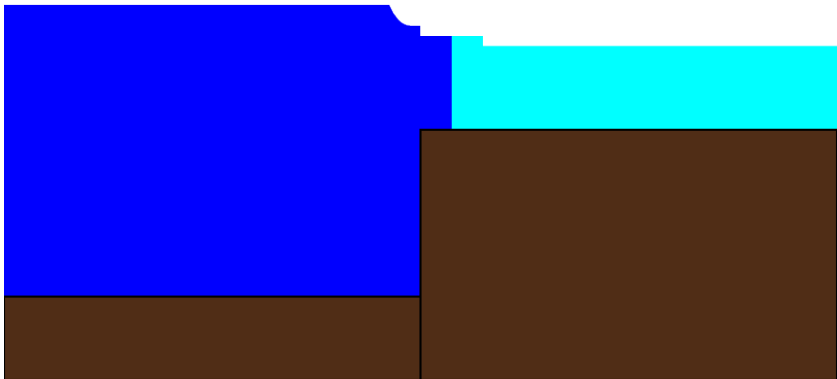
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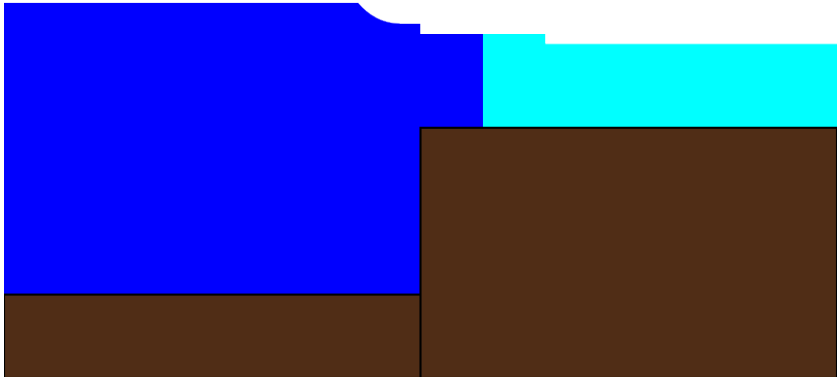
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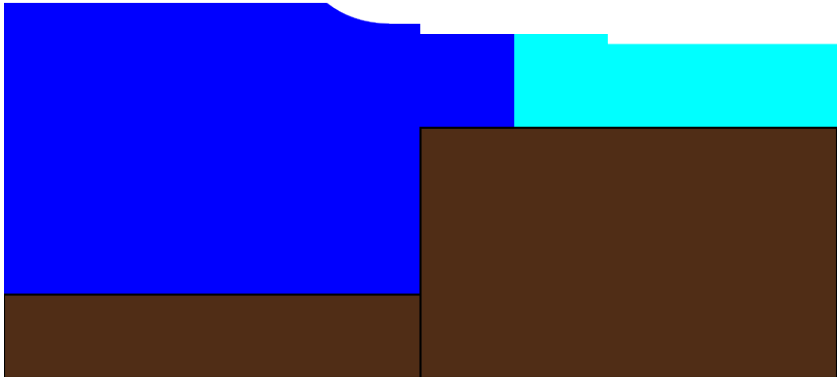
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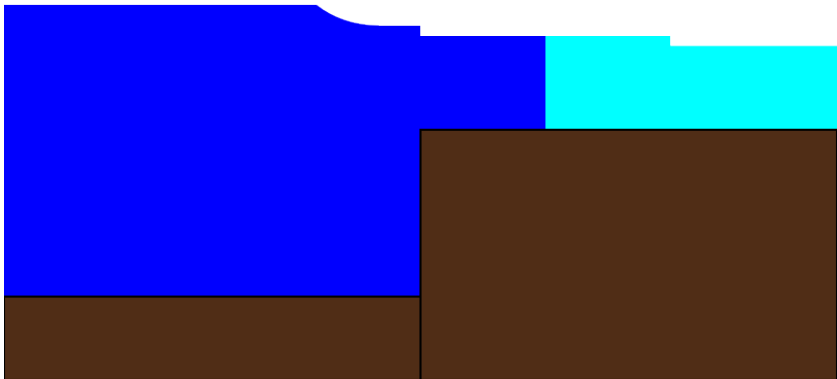
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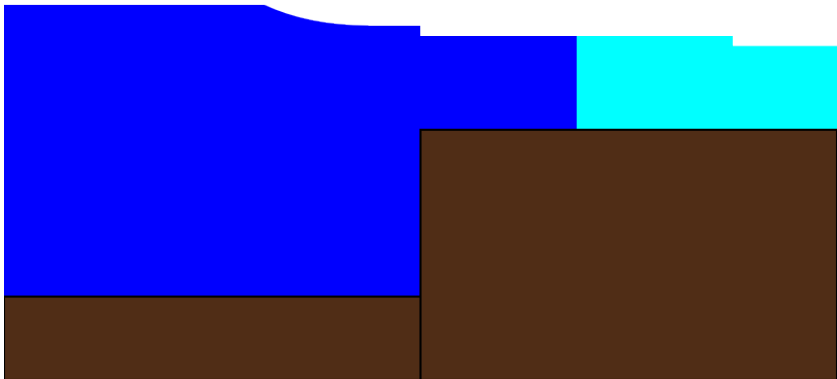
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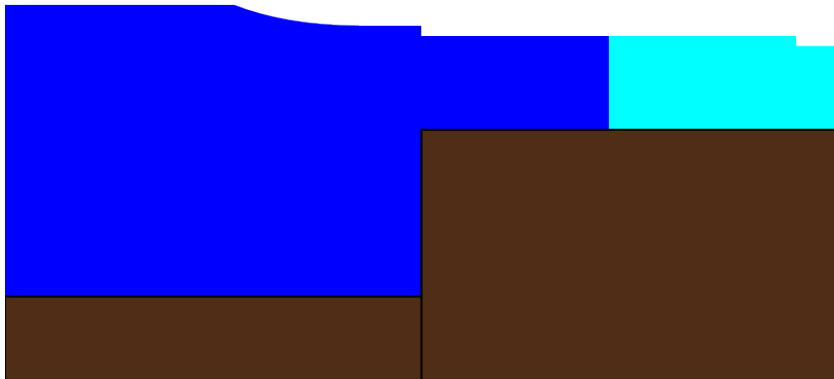
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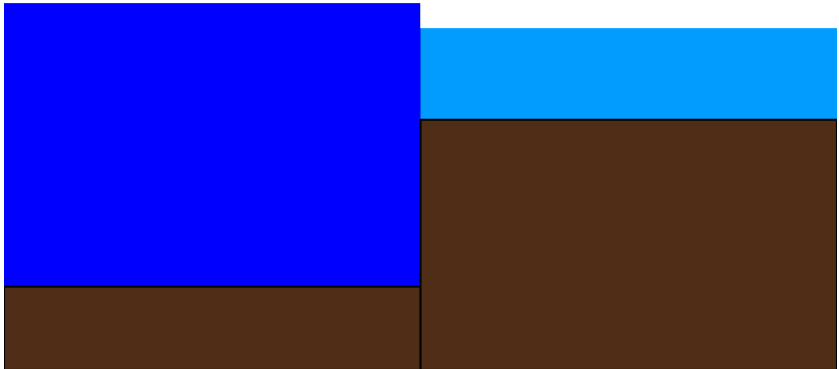
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Shorelines and inundation

- shallow water equations present difficulties where $h \rightarrow 0$
- determining the motion of the shoreline can be difficult
- often ad hoc approaches are used
- this is prone to numerical instabilities

$$\frac{\partial h}{\partial t} + \frac{\partial(hu)}{\partial x} = 0,$$
$$\frac{\partial(hu)}{\partial t} + \frac{\partial}{\partial x}\left(hu^2 + \frac{1}{2}gh^2\right) = -gh\frac{\partial b}{\partial x},$$

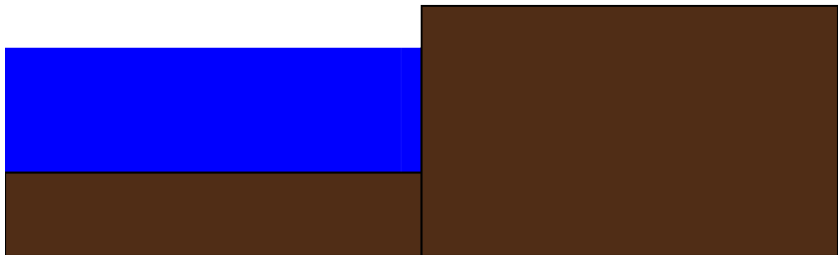
Shorelines and inundation

- the equations are only valid where $h > 0$
- the Riemann problem between wet and dry states can be exactly solved
- → only for the homogeneous problem (no bathymetry!)

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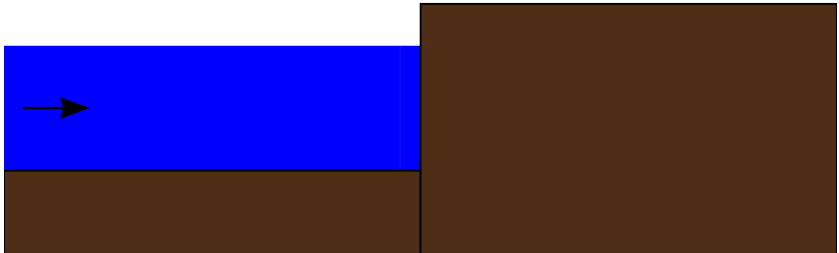
The Riemann problem at the shoreline.

motionless steady state at the shoreline: no waves.



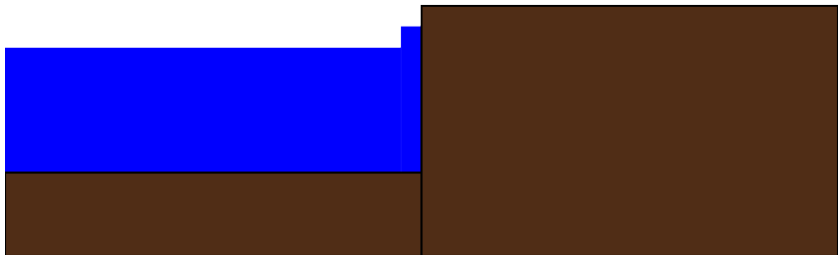
The Riemann problem at the shoreline.

flow at the shoreline: velocity insufficient for inundation.



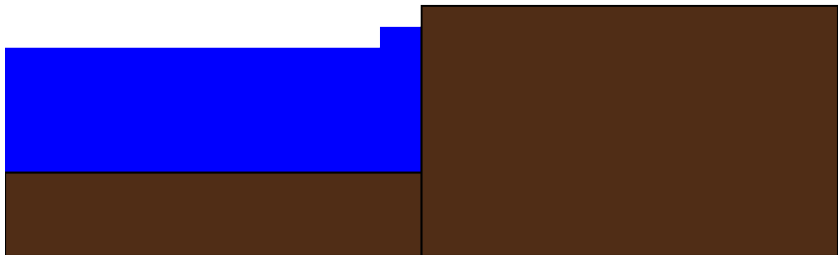
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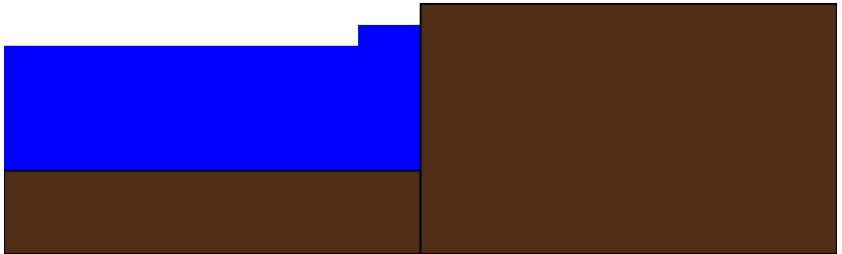
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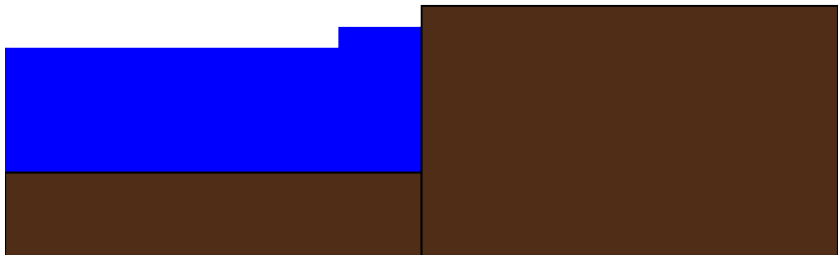
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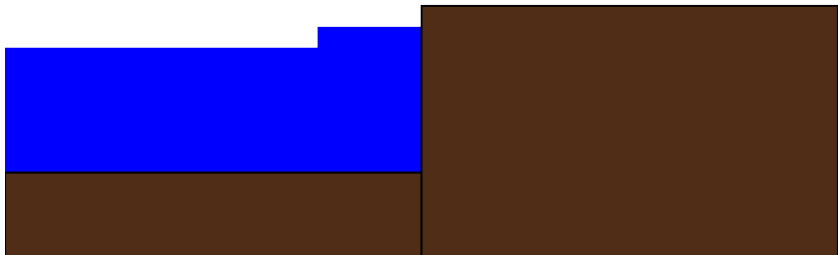
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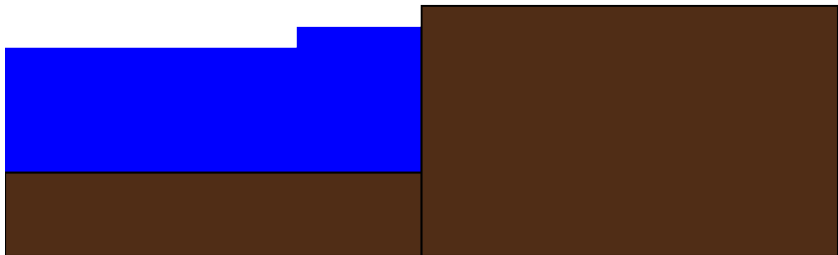
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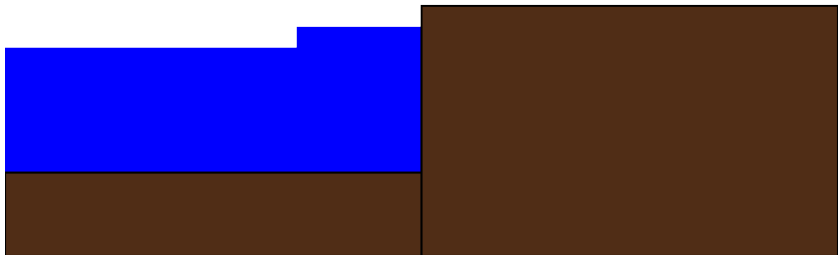
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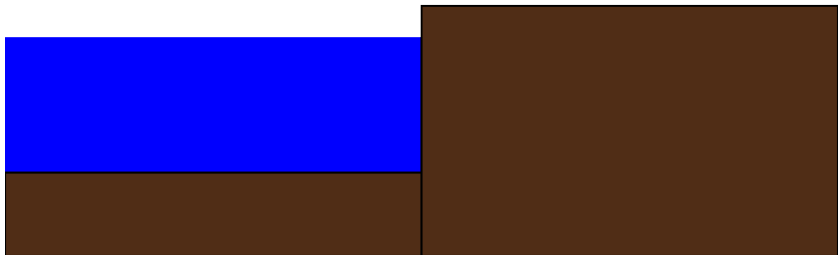
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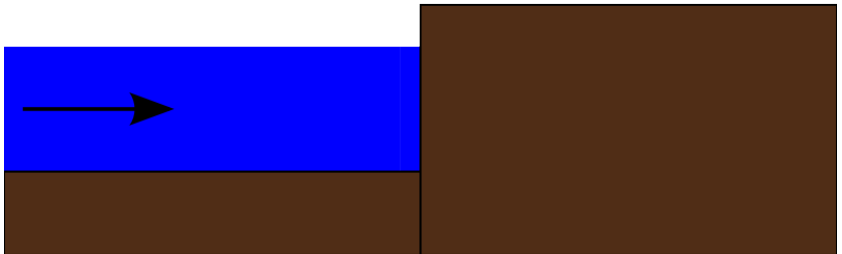
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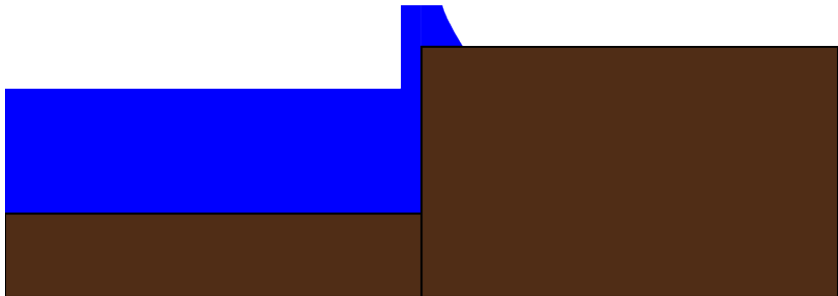
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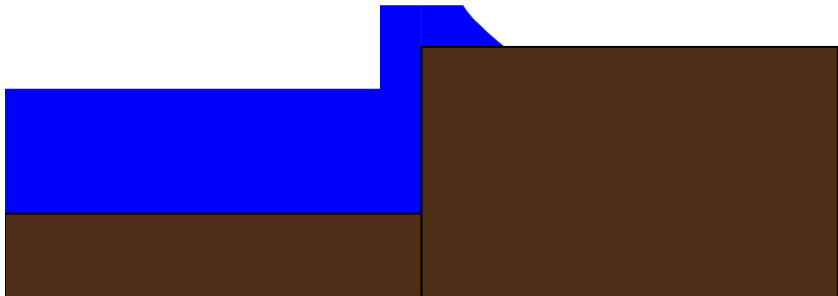
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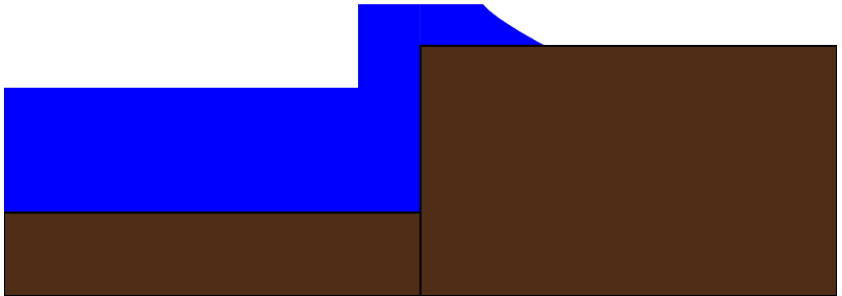
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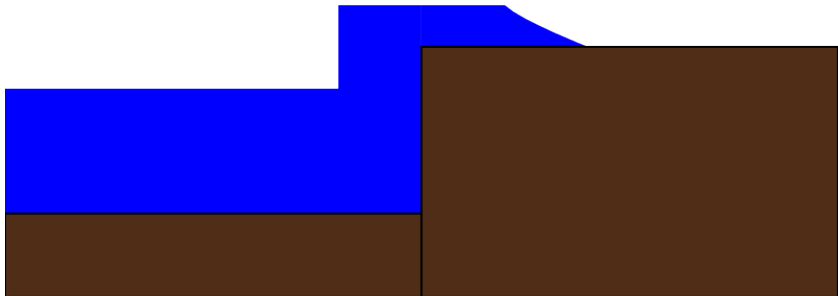
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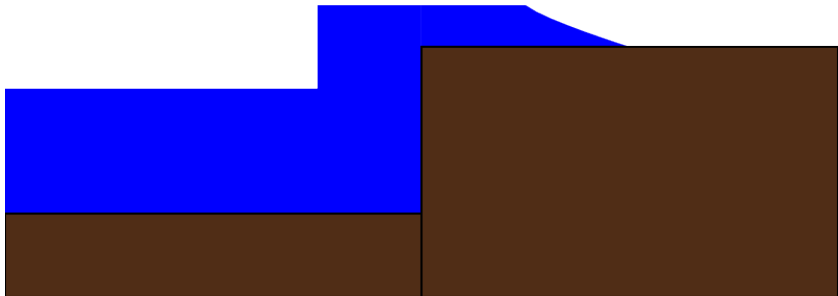
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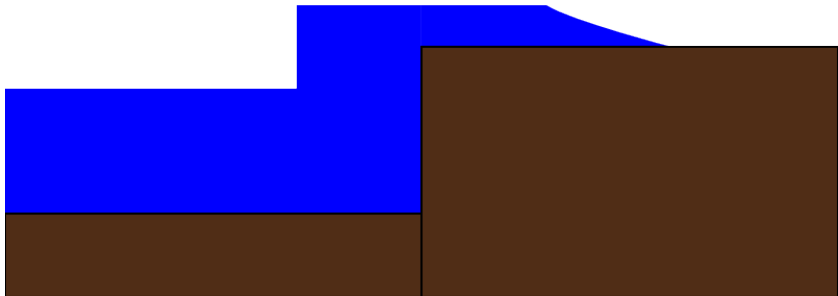
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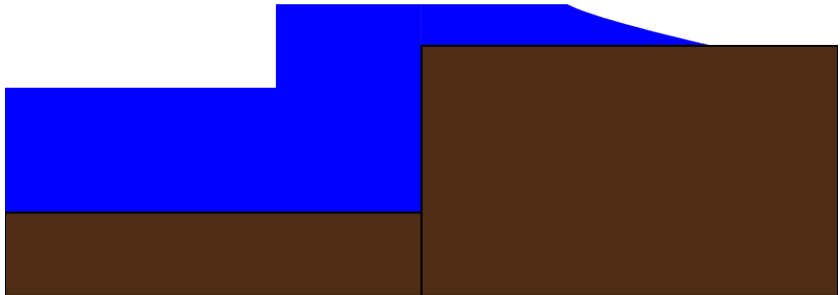
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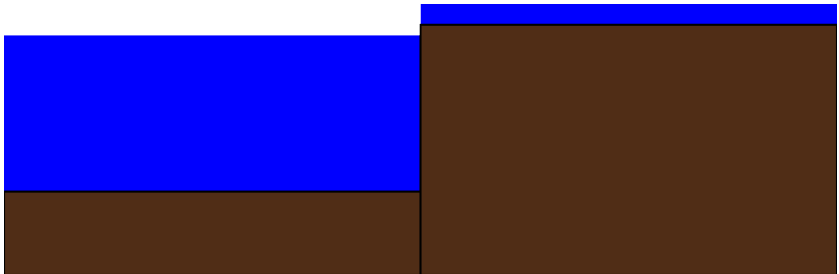
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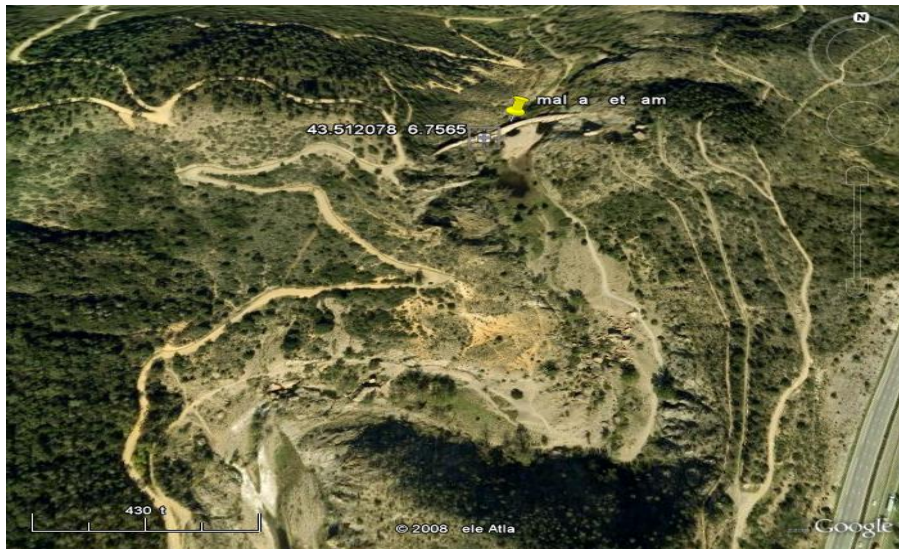


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flow at the shoreline: inundation.



Overland flooding: Malpasset dam, France 1959



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Landslides and Debris Flows

Debris flows, landslides etc.: liquified masses of soil and rock.



Indonesian Lahar Movie Ritigraben Switzerland Debris Flow Movie

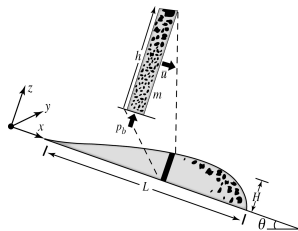
Debris-flow mobilization

Mobilization: $q_t + f(q)_x + \mathcal{W}(q)q_x = \psi(q, x)$

- failure occurs when the driving forces slightly exceed the shear strength at any one point: a small perturbation to a balanced steady state: $f(q)_x + \mathcal{W}(q)q_x \approx \psi(q, x)$
- failure: an equilibrium is perturbed...what happens next?
 - ① rapid temporary instability: (shear contraction \rightarrow increased pore pressure \rightarrow decreased shear strength \rightarrow landslide.)
 - ② quick stabilization: (shear dilation \rightarrow decreased pore pressure \rightarrow shear strength reestablished \rightarrow localized slump.)
 - ③ anything intermediate...e.g., stick-slip.
- modeling the outcome at least requires well-balanced methods...if it can be done at all. Numerical conservation must be maintained for $f(q)_x$ while simultaneously being well-balanced.

Mathematical model

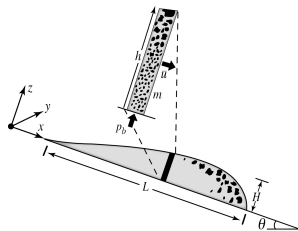
Model equations are depth-averaged:



- Integrating out the vertical component gives 2D system for h, u, v, m, p_b
- degree of accuracy lost upon depth-averaging depends on shallowness;
- computationally tractable for large-scale problems.

Mathematical model

Model equations are a hyperbolic system:



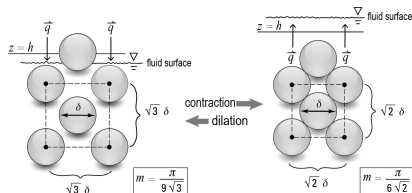
- similar in form to the shallow water equations;
- strictly hyperbolic system (desirable stability properties);

$$q_t + \mathcal{A}(q)q_x + \mathcal{B}(q)q_y = \psi(q),$$

$$\text{where } q = (h, hu, hv, hm, p_b)^T$$

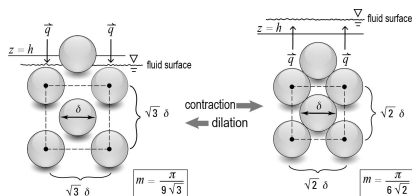
- coupled evolution of pore pressure and solid-volume fraction

Pore-fluid pressure solid-volume fraction coupling:



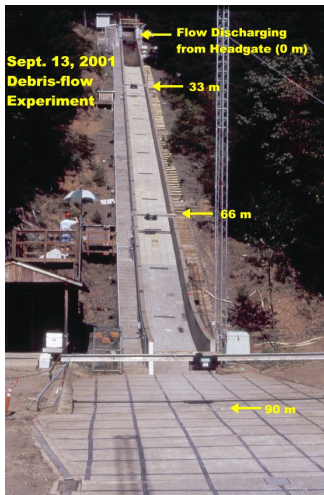
- dilation/contraction of solid phase affects pore-pressure
- pore-pressure mediates Coulomb solid stress through buoyancy (effective stress)
- frictional stress determines stability of sediment mass
- dilation/contraction based on dilatancy angle
- feedback loop: shearing \Leftrightarrow contraction/dilation \Leftrightarrow pore-pressure \Leftrightarrow frictional resistance

Modeling mobilization/stability:



- an initial rise in pore pressure can perturb stability
- evolution and feedback determine whether mass destabilizes into debris flow
- contractive case: shearing leads to higher pressure & less resistance (positive feedback)
- dilative case: shearing leads to lower pressure & higher resistance (negative feedback)

USGS experimental debris-flow flume



Play Movie

USGS experimental debris-flow flume

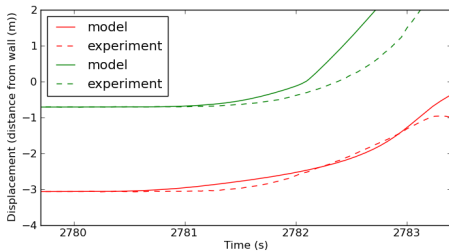
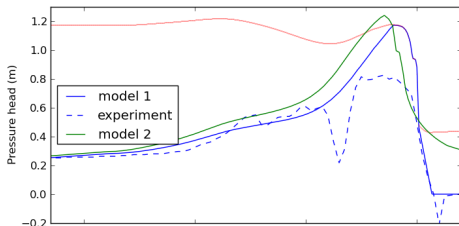
Simulating natural initiation by rising pore-pressure

USGS experimental debris-flow flume

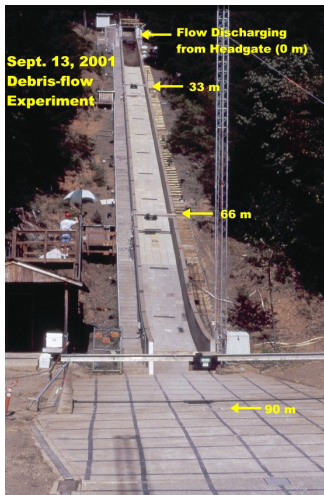
Simulating natural initiation by rising pore-pressure

USGS experimental debris-flow flume

Comparison of fluid-pore pressure



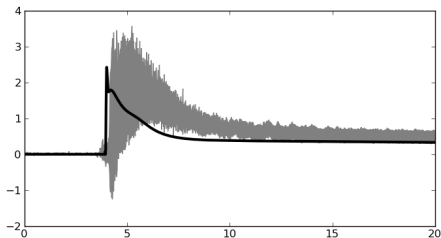
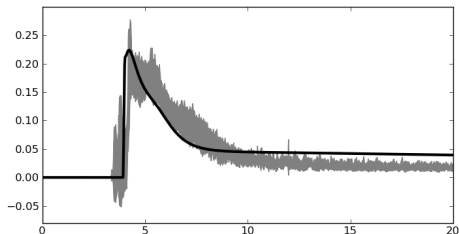
USGS experimental debris-flow flume



Play Movie

USGS experimental debris-flow flume

Comparison of downstream (32m) depth and fluid-pore pressure



Some future directions

- depth-averaged models can be coupled
 - landslide \rightarrow tsunamis through “dtopo”
 - landslides \leftrightarrow tsunami
 - granular-fluid model near inundation fronts
- always consider the possibility of numerical artifacts and/or instabilities in numerical solutions.
- validate, verify, compare, share.

Thank You!